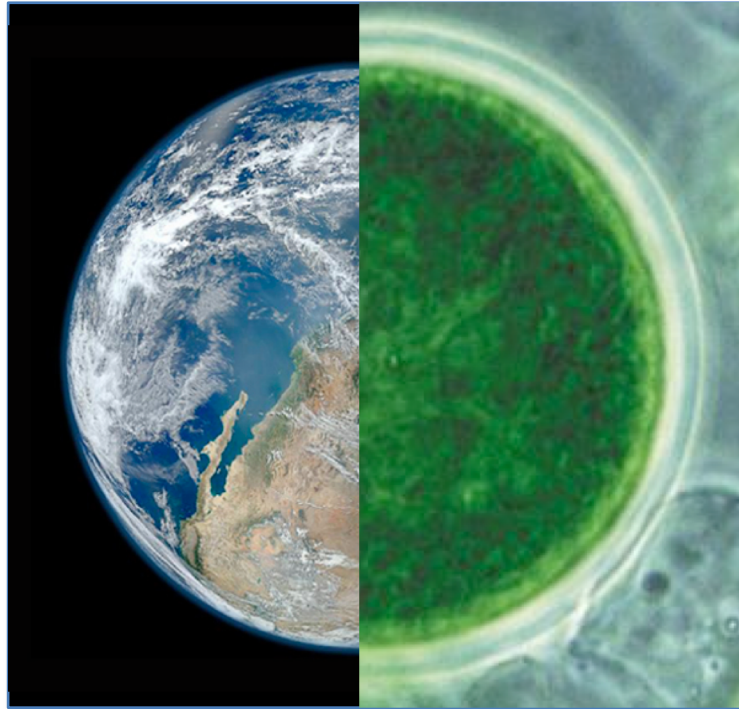


Co-evolution of life and environment on the early Earth



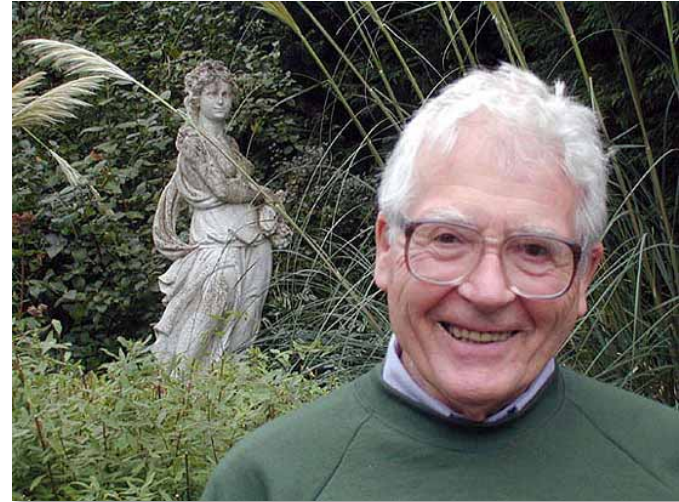
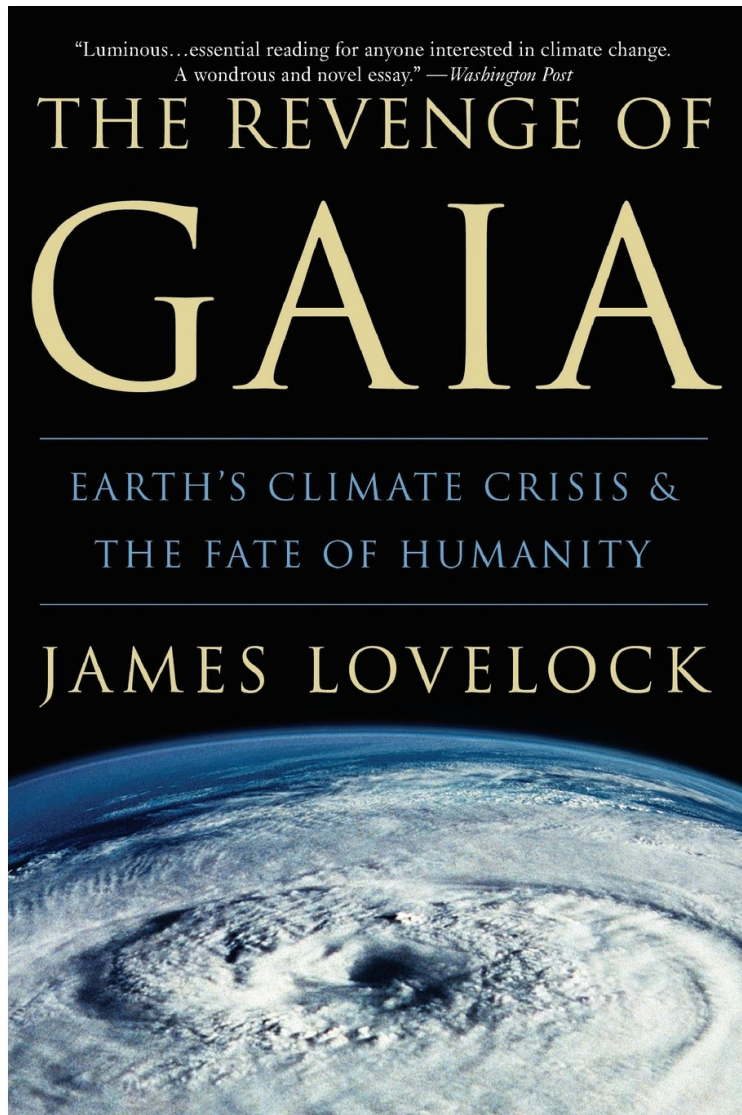
Benjamin Charnay¹, Boris Sauterey², Régis Ferrière²,
Franck Lefèvre³ et Stéphane Mazevet⁴

¹LESIA, Observatoire de Paris

²ENS Paris

³LATMOS, Paris

⁴LUTH, Observatoire de Paris

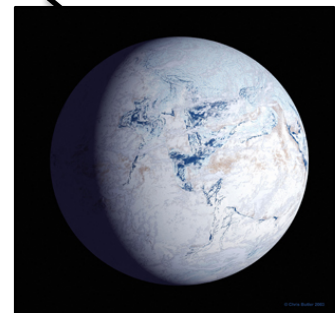
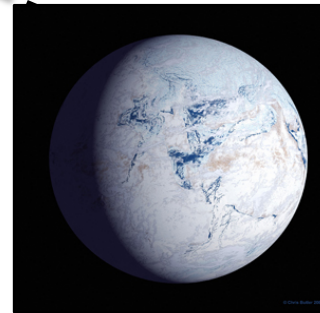
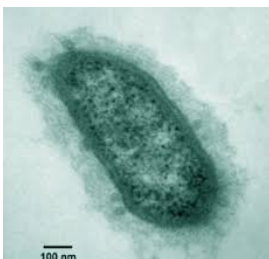
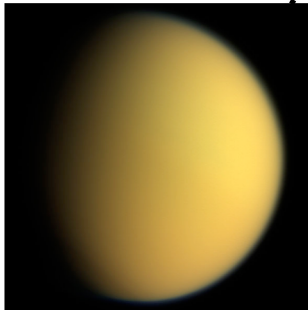
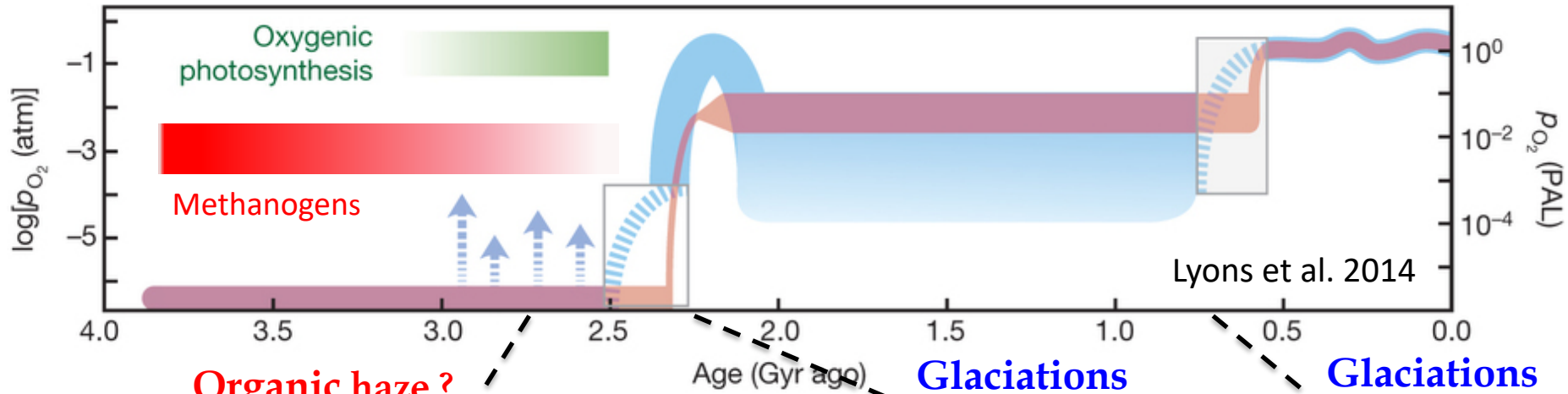


Gaia hypothesis:

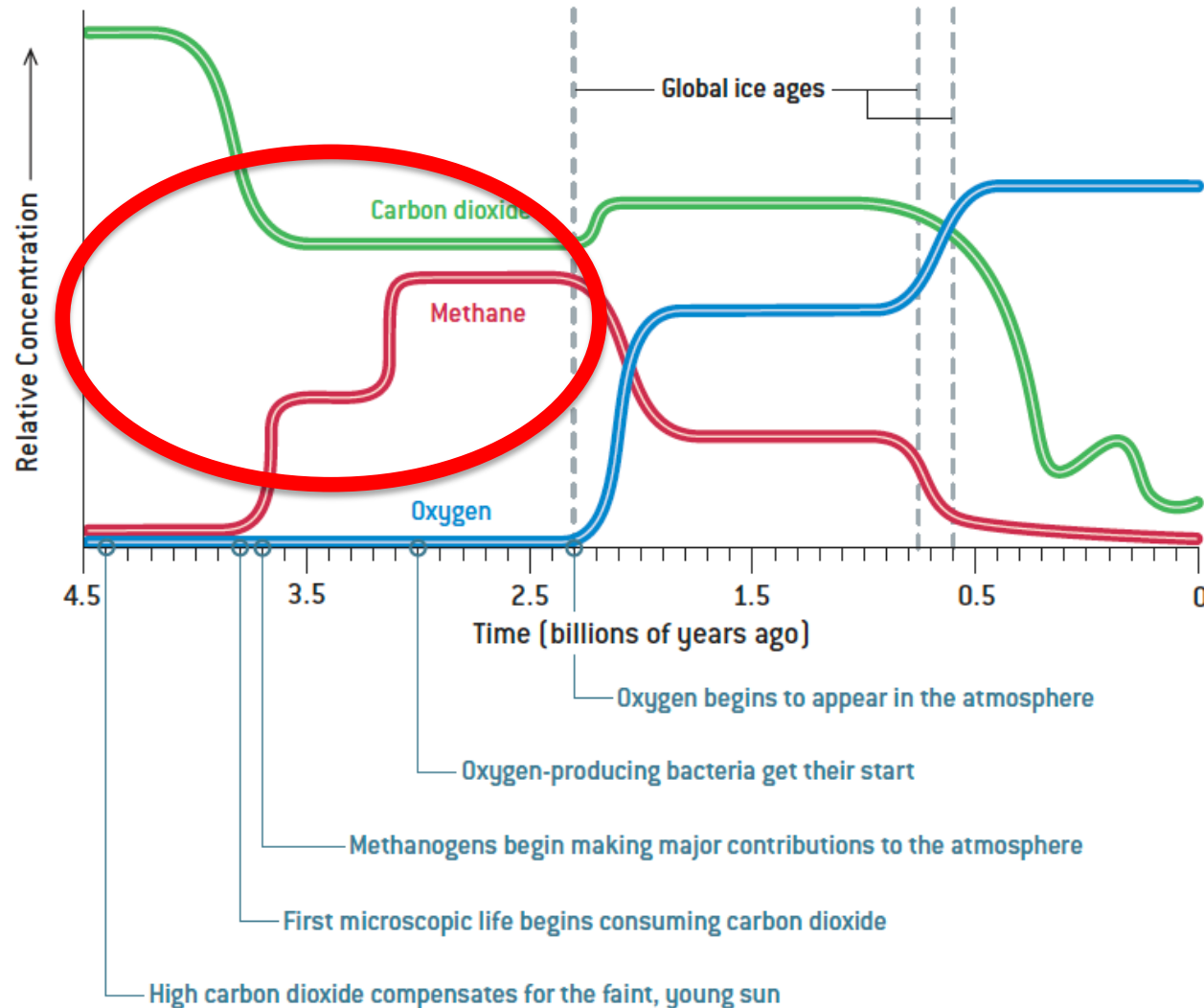
"A complex entity involving the Earth's biosphere, atmosphere, oceans, and soil; the totality constituting a feedback or cybernetic system which seeks an optimal physical and chemical environment for life on this planet."

Co-evolution of Earth's atmosphere and life

Evolution of O₂ in Earth's atmosphere



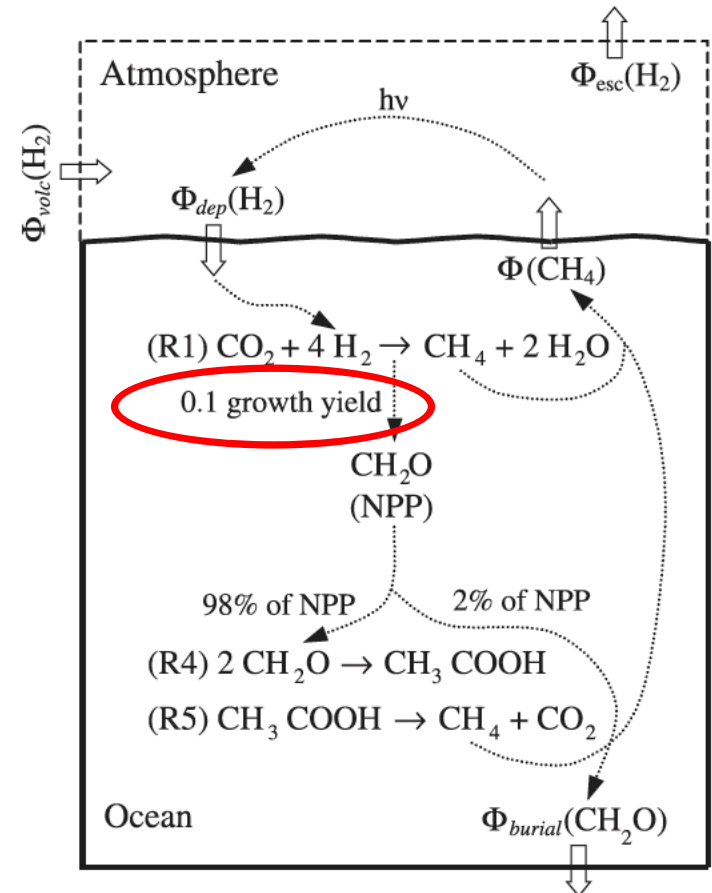
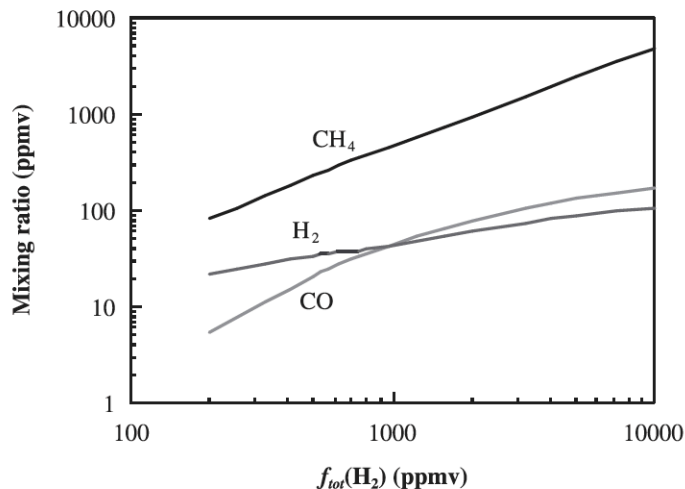
Co-evolution of Earth's atmosphere and life



Co-evolution of Earth's atmosphere and life

Previous study by [Karecha et al. 2005](#):

- Estimation of CH₄ production and NPP for different early ecosystems
- Calculation based on pure thermodynamics
- Fixed fraction of NPP and fixed ΔG



$$\Delta G = \Delta G^0 + RT \ln Q$$

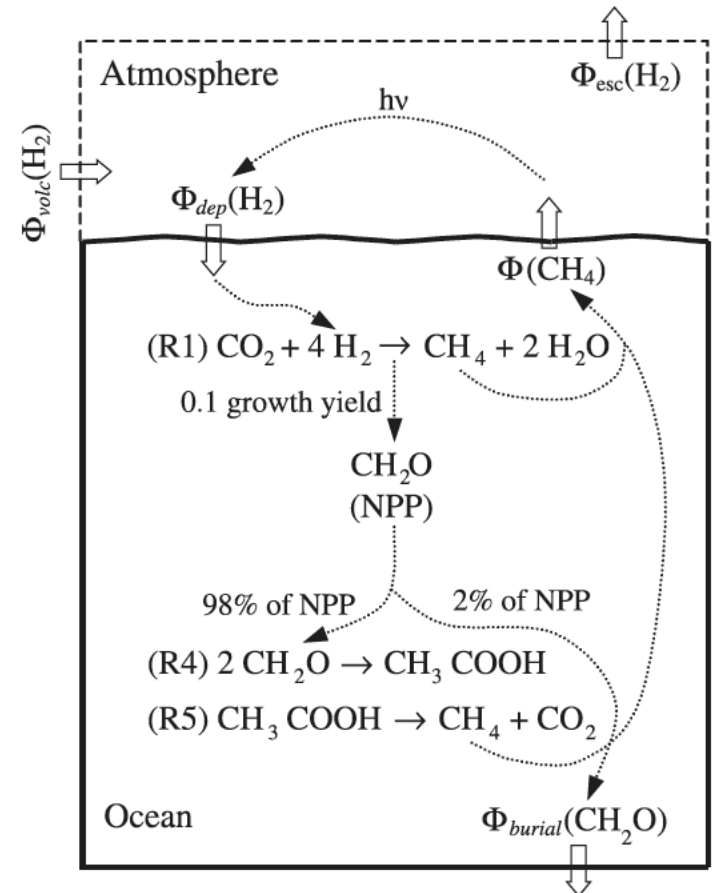
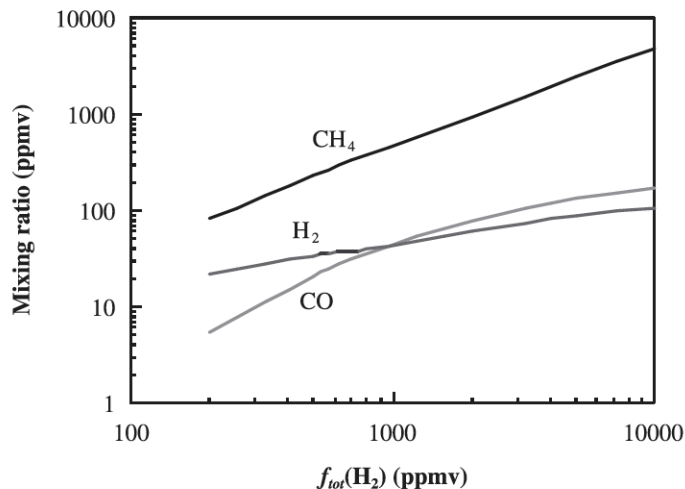
$$\Delta G = -30 \text{ KJ/mol (1 mol of ATP)}$$

$$Q = \frac{[CH_4]_{aq}^* a(H_2O)^2}{[CO_2]_{aq}^* ([H_2]_{aq}^*)^4}$$

Co-evolution of Earth's atmosphere and life

Previous study by [Karecha et al. 2005](#):

- Estimation of CH_4 production and NPP for different early ecosystems
- Calculation based on pure thermodynamics
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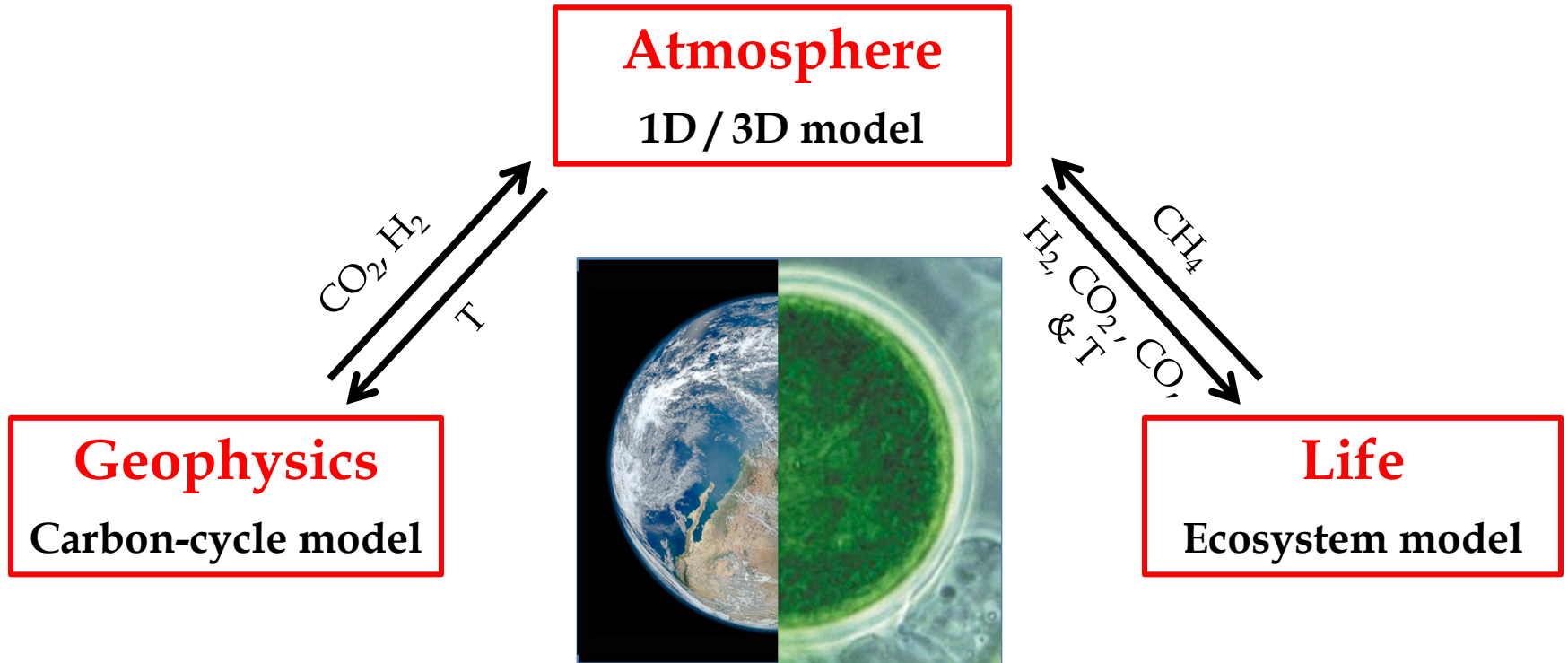


Goal of our study

Revisiting Karecha's study with state-of-the art models:

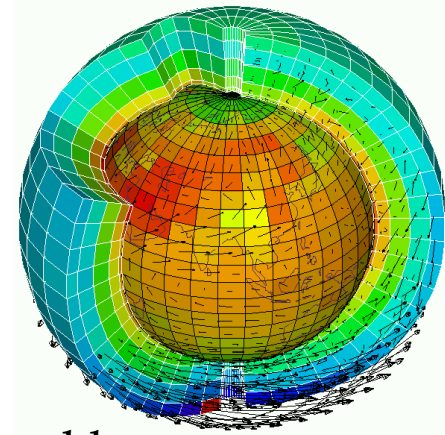
- Estimating CH_4 emission, greenhouse effect and NPP for different ecosystems
- Analysing feedbacks between early ecosystems and the environment

Co-evolution of Earth's atmosphere and life

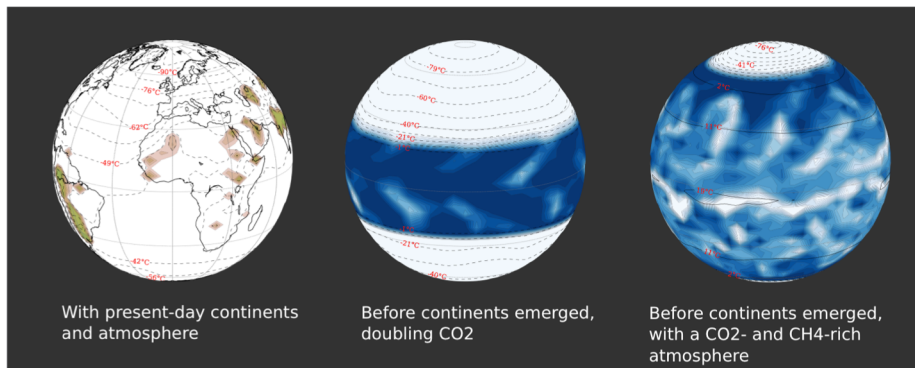


Atmospheric model:

- Climate states simulated with the **Generic LMD GCM**

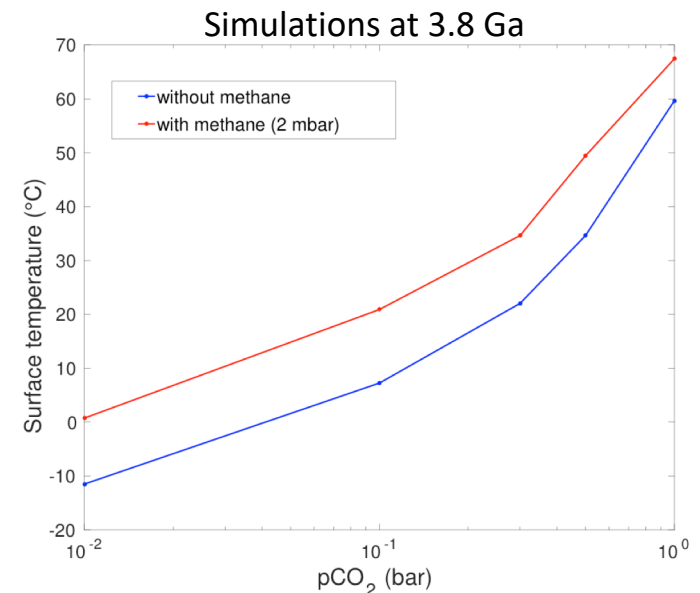


- Archean/Hadean climates and solutions to the faint young Sun problem (Charnay et al. 2013, 2018)



- Atmospheric composition: N₂, CO₂, CH₄ & H₂O
- Surface pressure = 1 bar
- No land

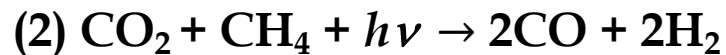
→ parameterization $T_{\text{surf}} = f(p\text{CO}_2, p\text{CH}_4, t)$



Atmospheric model:

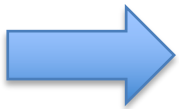
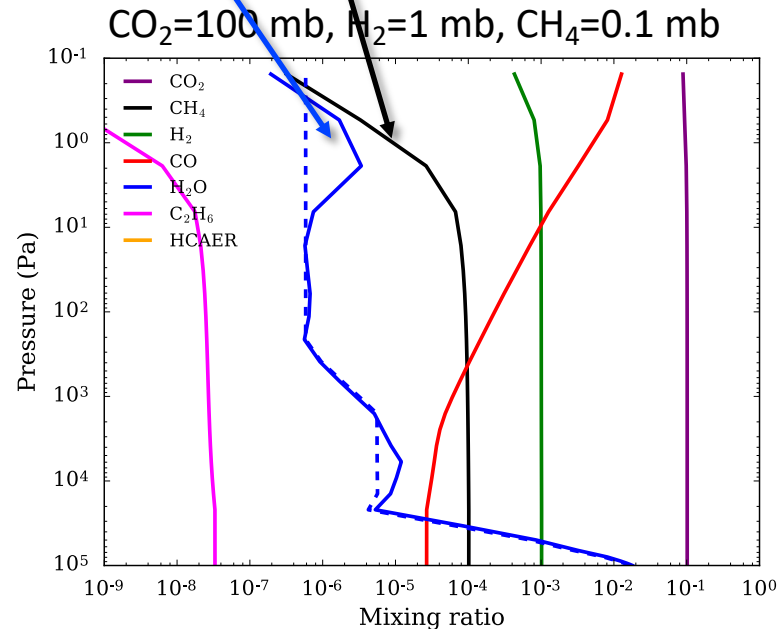
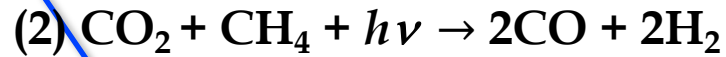
- Photochemistry in 1D with the photochemical core from [Lefèvre et al. 2003](#)
 - Hydrocarbure chemistry (18 species, 82 reactions)
 - Nitrogen chemistry (15 species, 42 reactions): Production of NO_x by lightning
 - Boundary conditions:
 - fixed mixing ratio (CO₂, H₂, CH₄)
 - fixed surface flux (CO, diffusion limited in the ocean)
 - H escape to space

Key reactions:



Atmospheric model:

Key reactions:

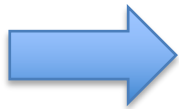
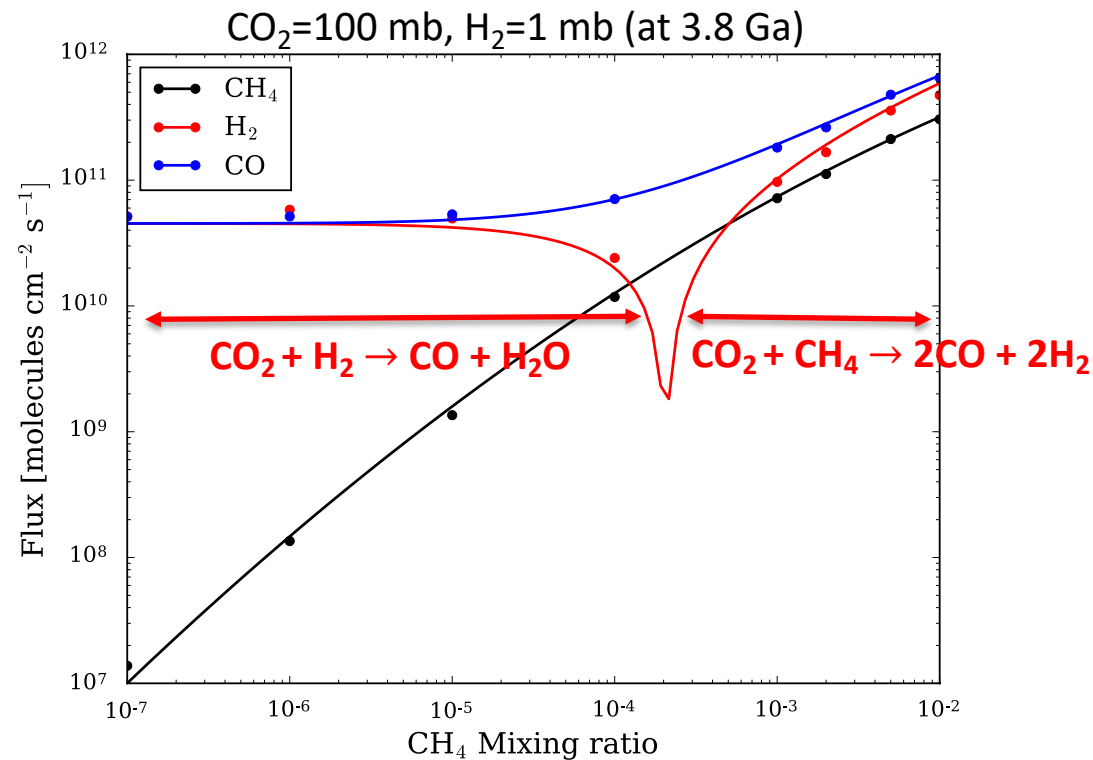


High CO production from CO_2 and CH_4 photolysis

Atmospheric model:

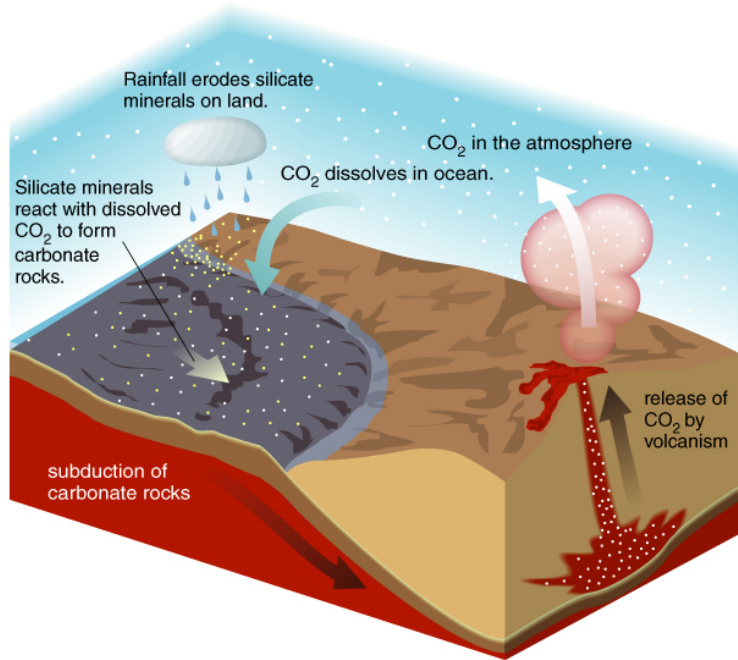
Computation of an atmospheric with different mixing ratio of CO_2 , H_2 and CH_4 (CO assumed to be efficiently consumed by acetogens)

→ parameterization for surface fluxes of CO_2 , H_2 , CH_4 and CO at equilibrium



High CO production from CO_2 and CH_4 photolysis

Carbon cycle model



© Addison-Wesley Longman

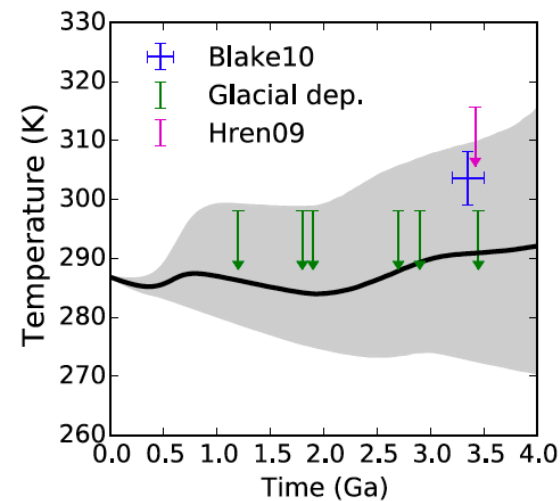
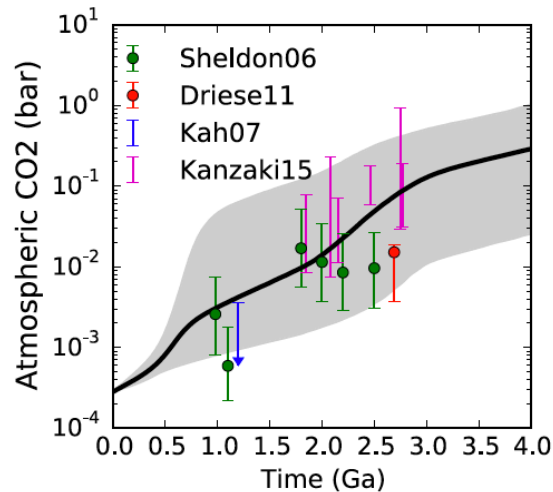
Model from Krissansen-Totton et al. 2018

CO_2 sources:

- Arc volcanoes
- Mid-oceanic ridges

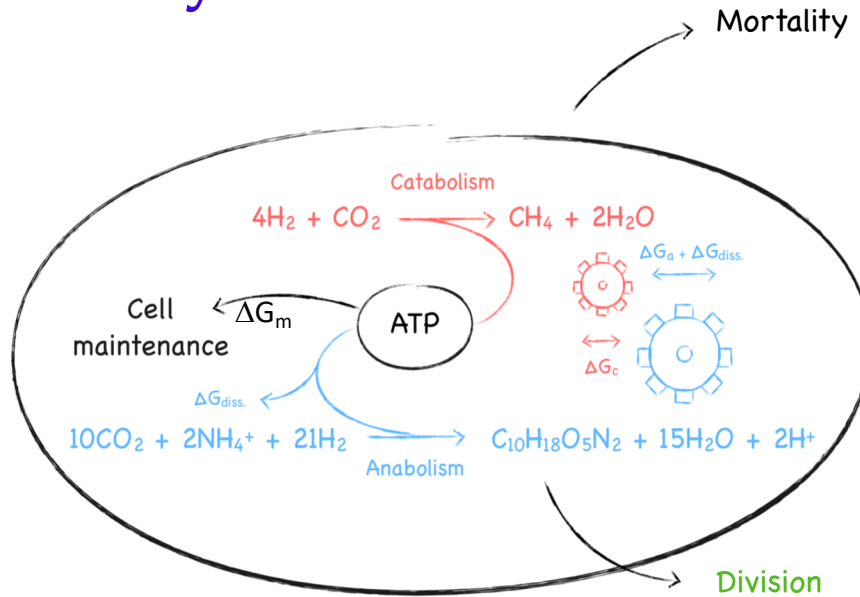
CO_2 sinks:

- Silicate weathering
- Seafloor weathering



Krissansen-Totton et al. 2018

Ecosystem model



Cellular dynamics :

- Catabolism
- Anabolism
- Cellular division
- Mortality

Population dynamics:

- Gas exchange with environment
- Influence of temperature

$$\frac{dN}{dt} = N \cdot (r(B) - m(q_{cat}))$$

$$\frac{dB}{dt} = q_{ana} - r(B) \frac{B + Q_c}{2}$$

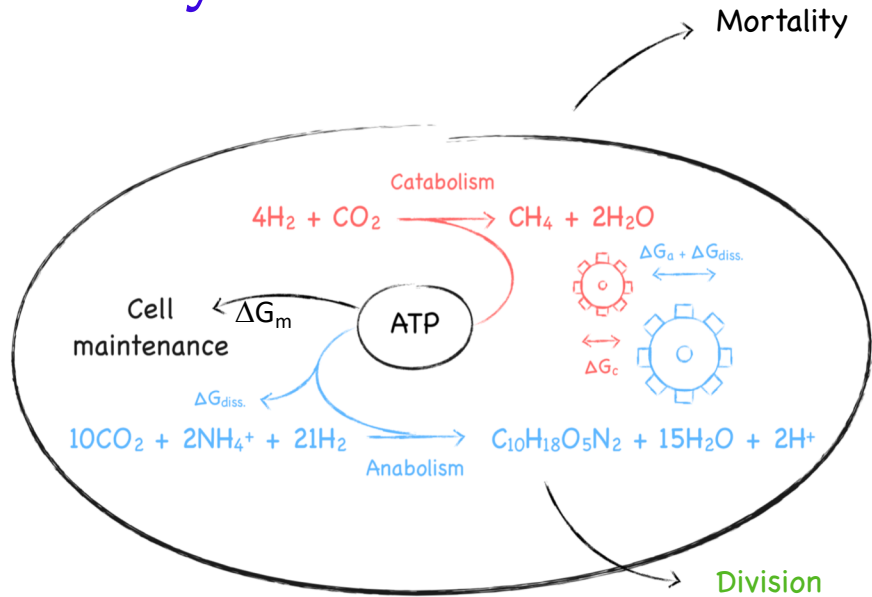
$$\begin{aligned} \frac{dX_i}{dt} = & F(X_i)_{bio}^{oc-atm} \\ & + N \cdot (q_{cat} \cdot \gamma_{Xi}^{cat} + q_{ana} \cdot \gamma_{Xi}^{ana}) \\ & + m \cdot \gamma_{Xi}^{rec} \cdot N(B + Q_c) \end{aligned}$$

N = cell concentration

B = biomass per cell

X_i = mixing ratio of specie i

Ecosystem model



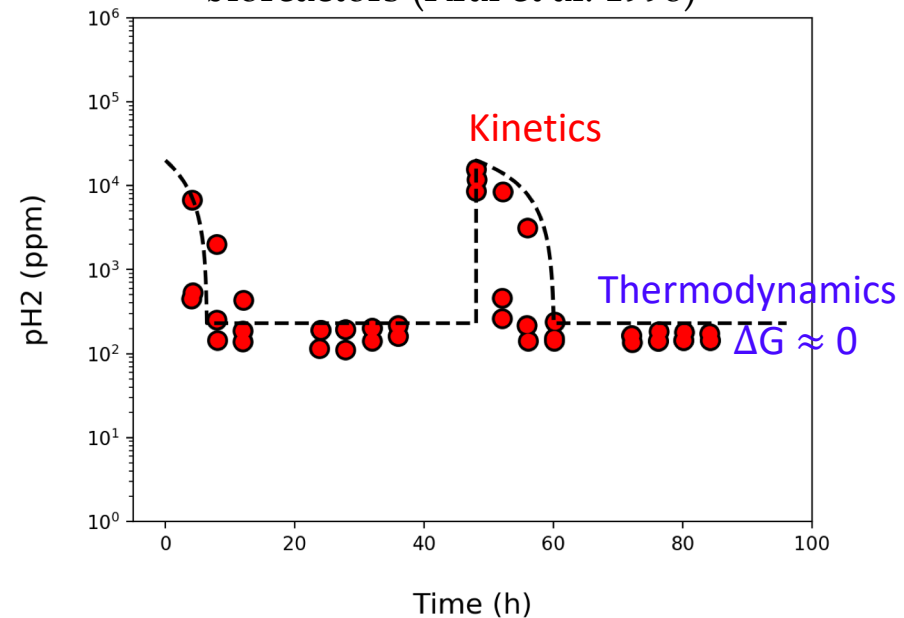
Cellular dynamics :

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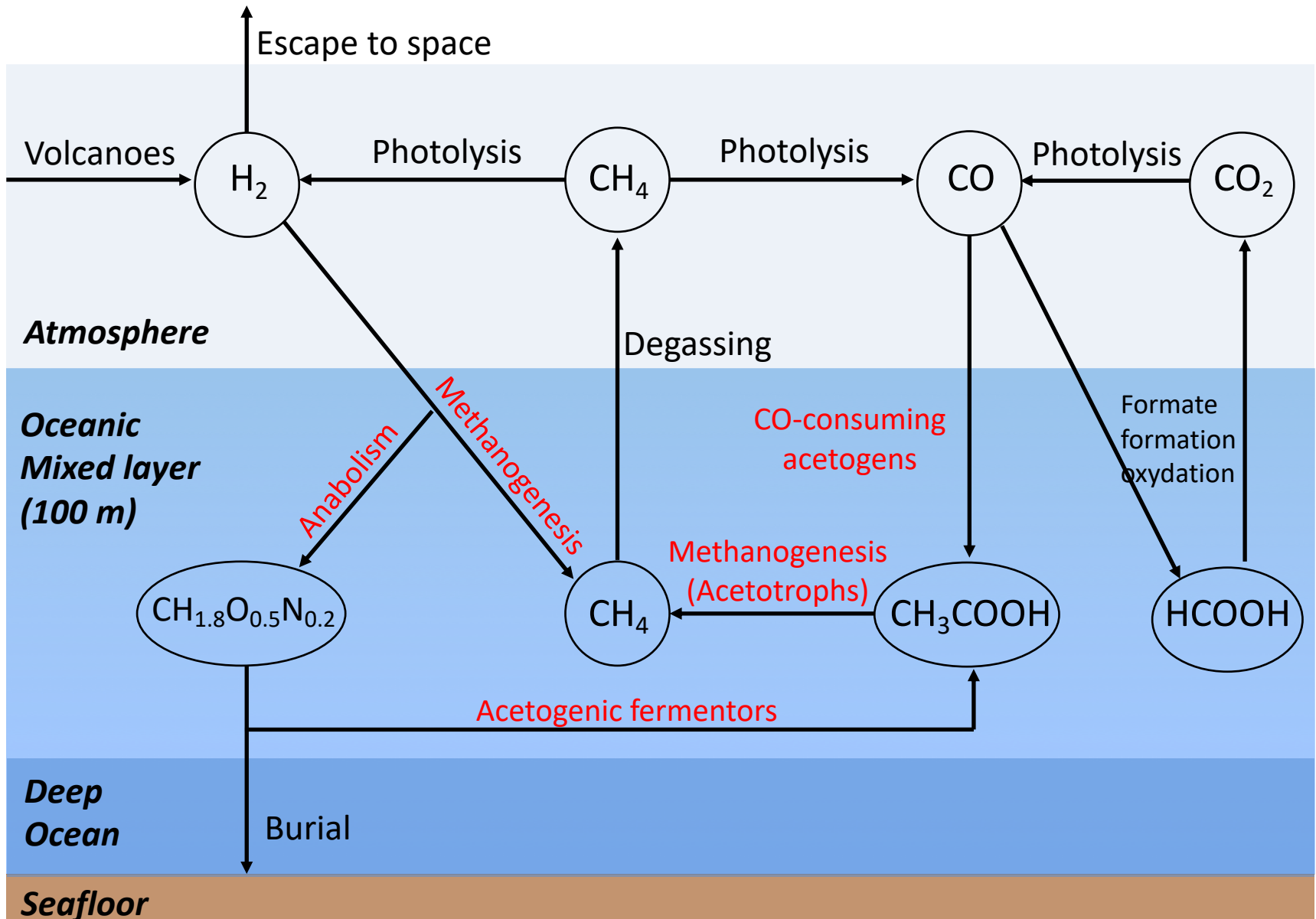
Validation with data of methanogens in bioreactors (Kral et al. 1998)



List of metabolisms

Reaction	ΔG_0	ΔH_0
Catabolic reactions:		
Methanogenesis: $0.25 \cdot CO_2 + H_2 \rightarrow 0.25 \cdot CH_4 + 0.5 \cdot H_2O$	-32.575	-63.175
Acetogenesis: $2 \cdot CO + H_2O \rightarrow CO_2 + 0.5 \cdot CH_3COOH$	-77.850	-129.850
Acetotrophy: $CH_3COOH \rightarrow CO_2 + CH_4$	-55.0	16.2
Acetogenic fermentors: $CH_{1.8}O_{0.5}N_{0.2} + 5/6 H_2O + 0.2 \cdot H^+ \rightarrow 1/3 CH_3COOH + 1/3 CO_2 + 2.3/3 H_2 + 0.2 \cdot NH_4^+$	-12.71	10.066
Anabolic reactions:		
$CO_2 + 0.2 \cdot NH_4^+ + 2.1 H_2 \rightarrow CH_{1.8}O_{0.5}N_{0.2} + 1.5 \cdot H_2O + 0.2 H^+$	-12.390	-99.700
$CO_2 + 0.1 \cdot N_2 + 2.1 H_2 \rightarrow CH_{1.8}O_{0.5}N_{0.2} + 1.5 \cdot H_2O$	28.25	128

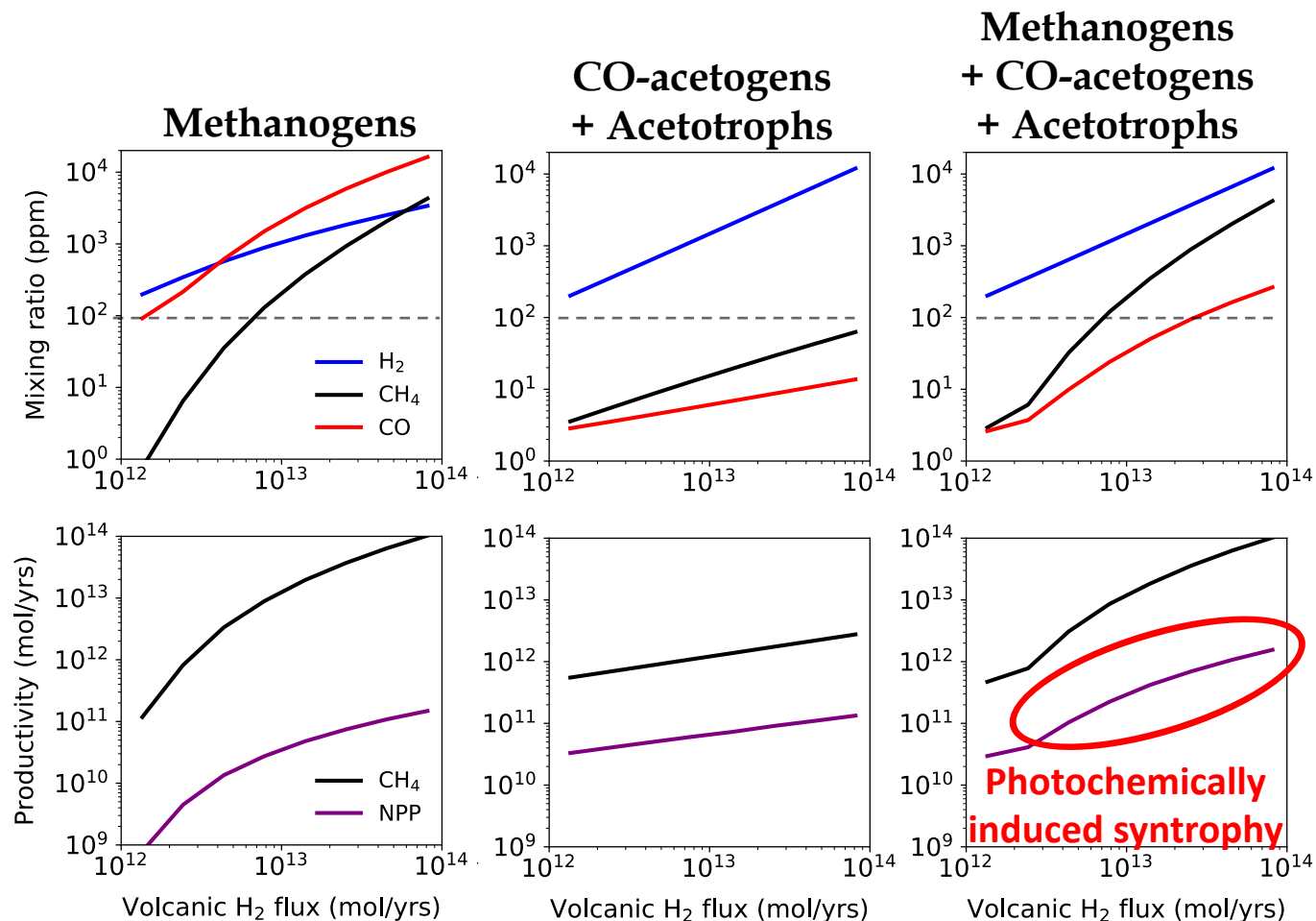
Diagram of the primitive biosphere



Results for ecosystems + atmosphere

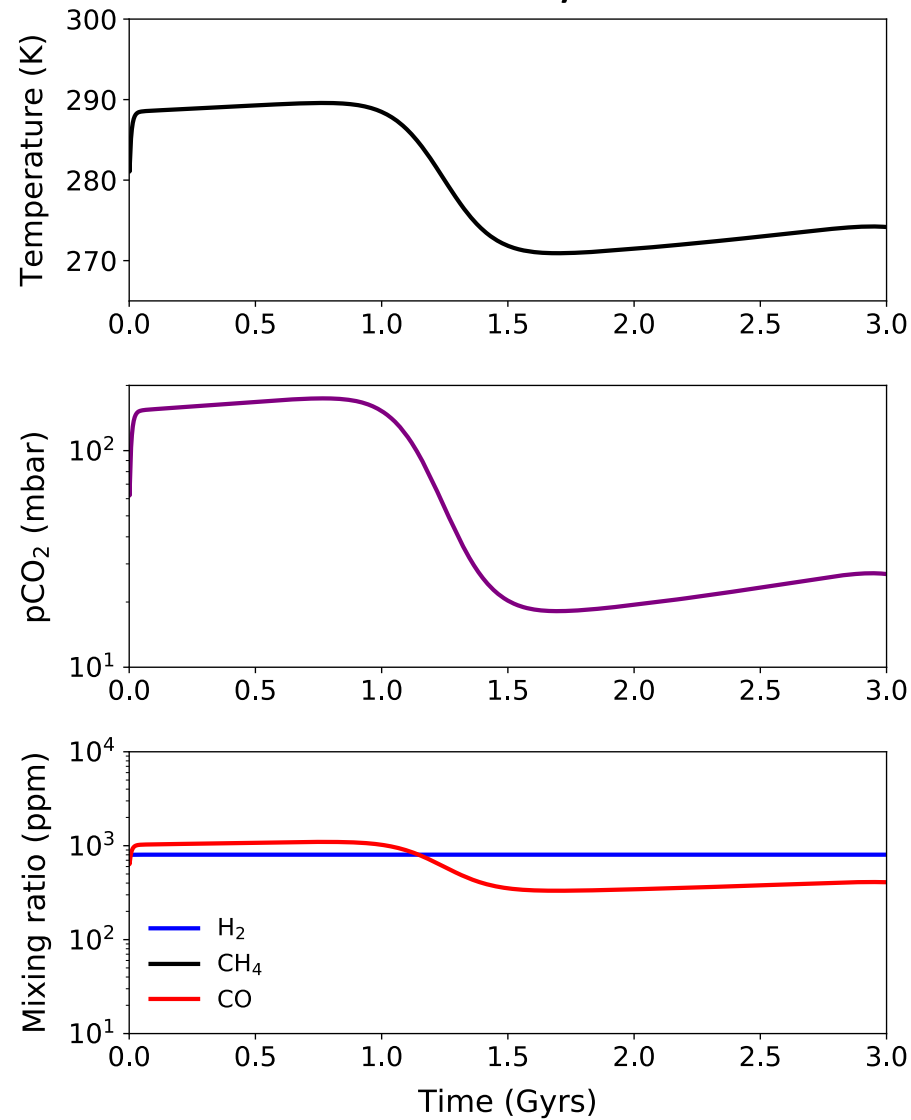
Simulations at 3.8 Ga with :

- $p\text{CO}_2 = 100 \text{ mbar}$, $T = 285 \text{ K}$
- $\Phi(\text{H}_2)_{\text{volc}} = 1\text{-}80 \text{ Tmol / yr}$
- non-limited by nitrogen (N-fixation or recycling by fermentors)



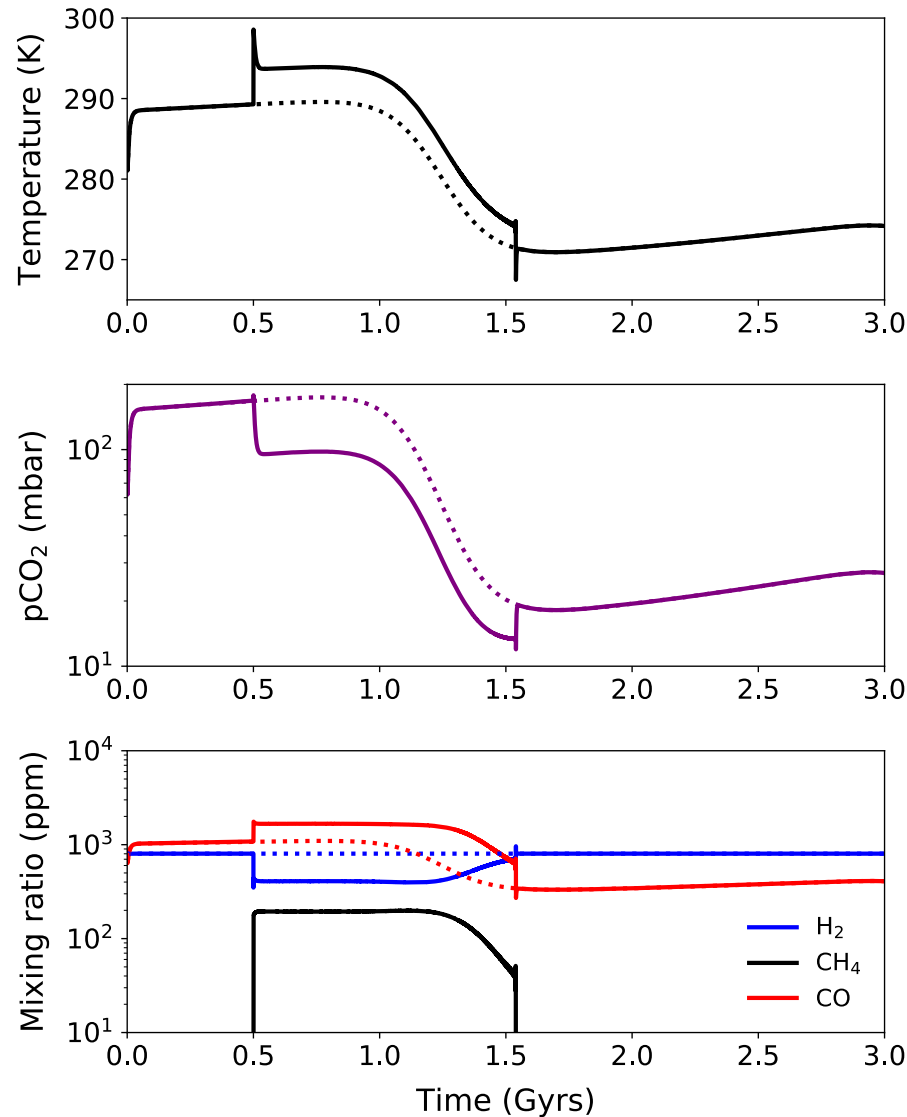
Results for ecosystems + atmosphere + C cycle

Without ecosystem



Results for ecosystems + atmosphere + C cycle

With methanogens after 0.5 Gyr



Possible implications for the triggering of Huronian glaciation

Summary

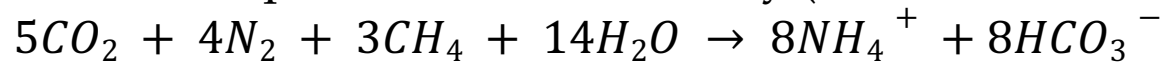
- First dynamic model of early ecosystems coupled to climate and C cycle
- Efficient CH₄ production by methanogens ([CH₄] = 100-1000 ppm) but less than Karecha et al. 2005
- Low NPP by methanogens (~30× lower than Karecha)
- CO-acetogens + acetotrophs enhance NPP by a factor ~10
- Methanogens induce a weak positive feedback on climate and facilitate glaciations



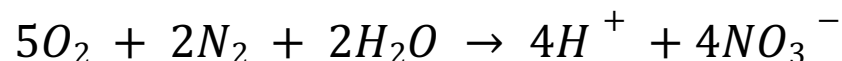
No gaian regulation of the climate by methanogens

Perspectives

- Including anoxygenic and oxygenic photosynthesis
- Computation of chemical disequilibrium and biosignatures
For the Archean, disequilibrium dominated by (Krissansen-Totton et al. 2018):

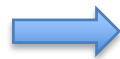
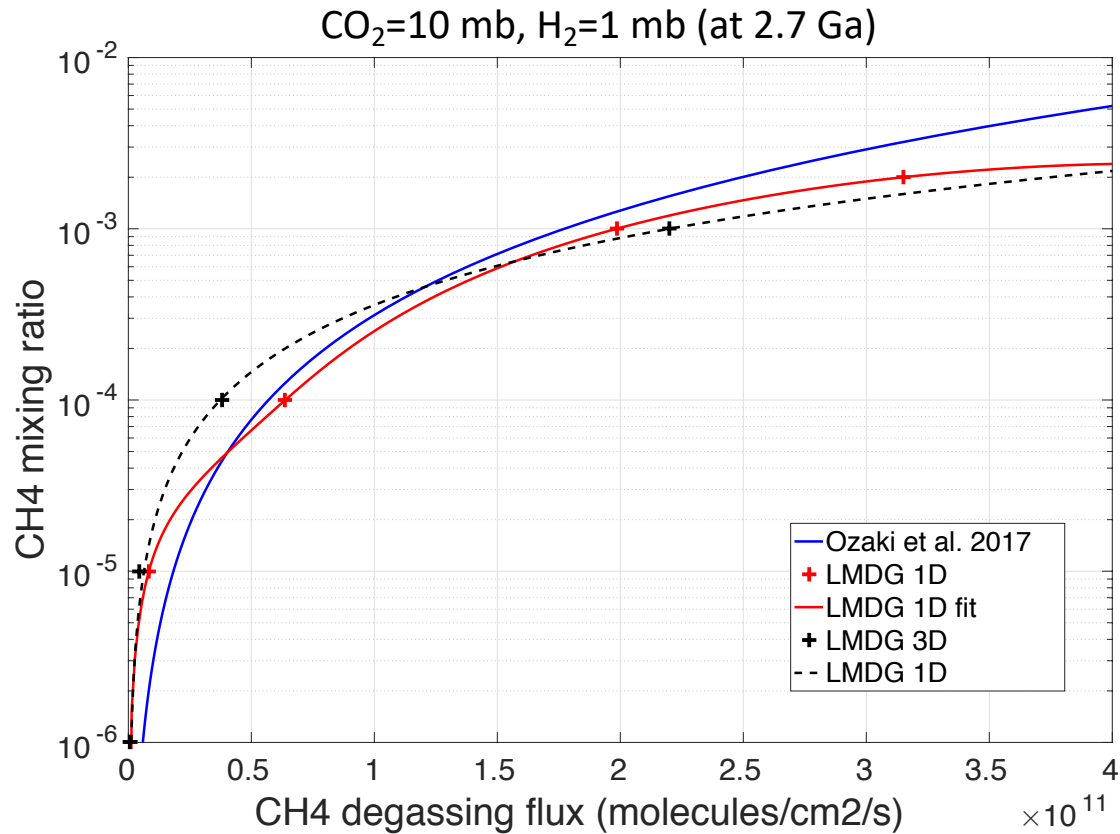


For the Proterozoic and modern Earth:



Atmospheric model:

Comparison to previous study and 1D VS 3D

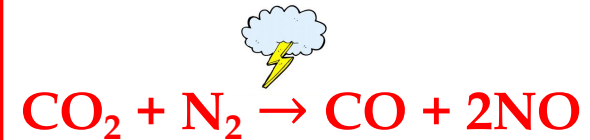


1D model adequate

Atmospheric model:

- Nitrogen chemistry (15 species, 42 reactions)

Production of NO_x using lightning frequency from [Wong, Charnay et al. 2017](#)



At 3.8 Ga with $p\text{CO}_2=0.1$ bar:

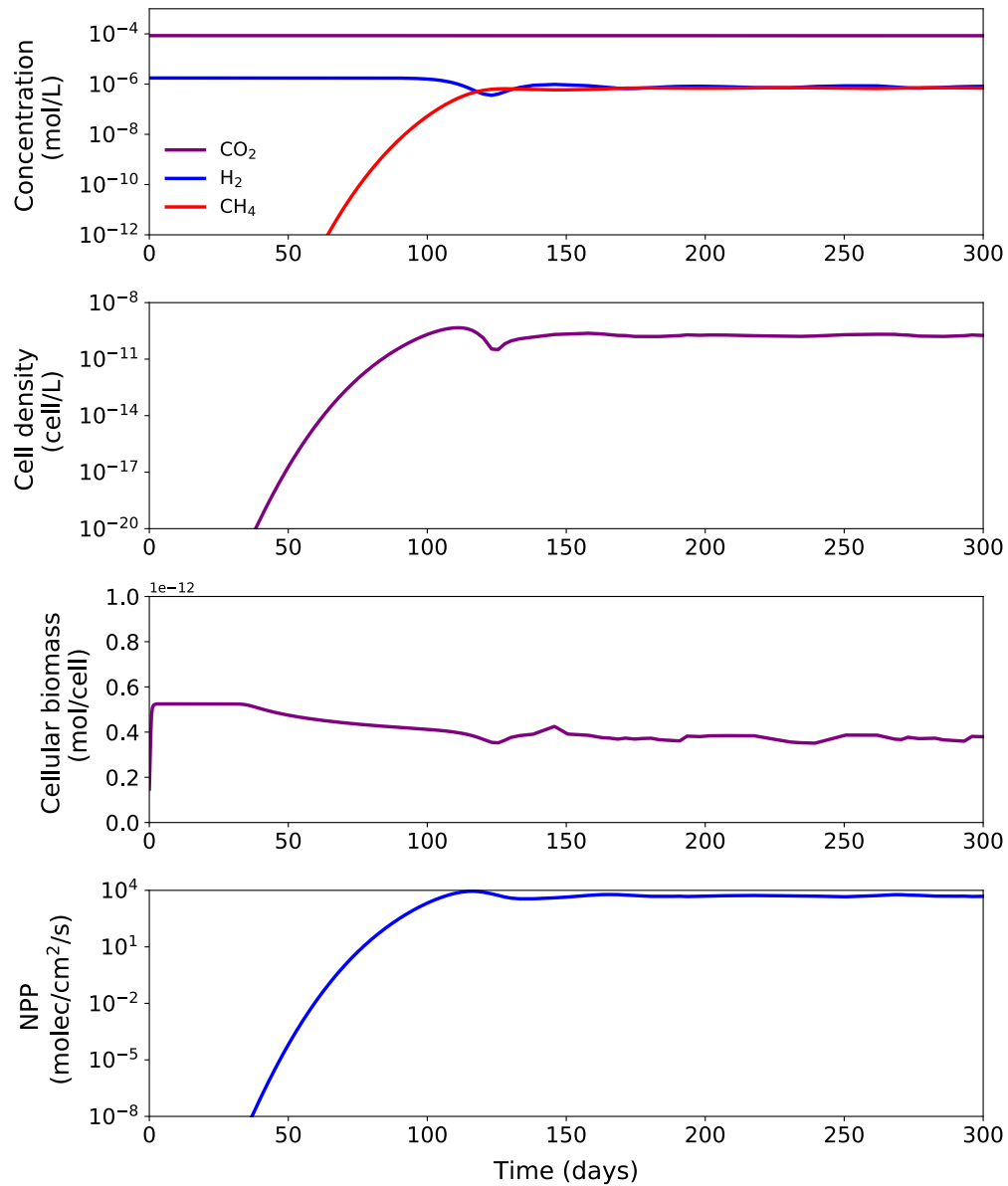
NO_x flux in the ocean = 1.8×10^8 mol/yr (essentially as HNO)

List of metabolisms

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Acetogenesis: $2 \cdot \text{CO} + \text{H}_2\text{O} \rightarrow \text{CO}_2 + 0.5 \cdot \text{CH}_3\text{COOH}$	-77.850	-129.850
Acetotrophy: $\text{CH}_3\text{COOH} \rightarrow \text{CO}_2 + \text{CH}_4$	-55.0	16.2
NO₃ Methanotrophy: $0.25 \cdot \text{CH}_4 + \text{NO}_3^- \rightarrow 0.25 \cdot \text{CO}_2 + \text{NO}_2^- + 0.5 \cdot \text{H}_2\text{O}$	-125.5	-120
NO₂ Methanotrophy: $0.375 \cdot \text{CH}_4 + \text{NO}_2^- + \text{H}^+ \rightarrow 0.375 \cdot \text{CO}_2 + 0.5 \cdot \text{N}_2 + 1.25 \cdot \text{H}_2\text{O}$	-393.14	-372.24
<div> <div></div> <div>low flux of NOx</div> </div>		
Acetogenic fermentors: $\text{CH}_{1.8}\text{O}_{0.5}\text{N}_{0.2} + 5/6 \text{H}_2\text{O} + 0.2 \cdot \text{H}^+ \rightarrow 1/3 \text{CH}_3\text{COOH} + 1/3 \text{CO}_2 + 2.3/3 \text{H}_2 + 0.2 \cdot \text{NH}_4^+$	-12.71	10.066
Anabolic reactions:		
$\text{CO}_2 + 0.2 \cdot \text{NH}_4^+ + 2.1 \text{H}_2 \rightarrow \text{CH}_{1.8}\text{O}_{0.5}\text{N}_{0.2} + 1.5 \cdot \text{H}_2\text{O} + 0.2 \text{H}^+$	-12.390	-99.700
$\text{CO}_2 + 0.1 \cdot \text{N}_2 + 2.1 \text{H}_2 \rightarrow \text{CH}_{1.8}\text{O}_{0.5}\text{N}_{0.2} + 1.5 \cdot \text{H}_2\text{O}$	28.25	128

Dynamics of methanogens

($p\text{CO}_2 = 100 \text{ mbar}$, $p\text{H}_2 = 1 \text{ mbar}$, $T = 285 \text{ K}$)



Present day Earth under a 20% weaker Sun

