



UNIVERSITÉ DE STRASBOURG



Université

de Strasbourg

Initiative d'excellence

La Géobiotropie

ou l'eau dans l'état en dessous du point critique pour la formation des minéraux de fer ferrique et des molécules de la vie

extraits des articles :

Bassez MP, Orig Life Evol Biosph 47:453-480 (2017)

Bassez MP, Orig Life Evol Biosph (accès libre et en ligne le 8 August 2018)

Bassez MP, en preparation (2018)

Marie-Paule BASSEZ

<http://chemphys.u-strasbg.fr/mpb>

Rencontres de la Société Française d'Exobiologie, S^t Michel l'Observatoire, 15-18 Oct 2018

Four processes all occurring in high-subcritical water, hscw lead to the formation of ferric minerals and molecules of life

critical point of water: $T_c=374^\circ\text{C}$, $P_c=22.1\text{MPa}$

hsc water: $300^\circ\text{-}350^\circ\text{C}$, $10\text{-}25\text{ MPa}$,
 $\text{pH} \sim 9.5\text{-}14$

1. **Oxidation** of Fe^{II} into Fe^{III}
concomitant to reduction of the H^+ of water to **produce H_2**
2. High **Solubility** of SiO_2 , **quartz & amorphous silica**
3. **Interaction** between **Fe^{II} -silicates** and hscw undersaturated in silica
for production of Fe^{III} -oxides & Fe^{III} -silicates (Banded Iron Formations) and H_2
4. **Hydrogenation** of CO_2 by H_2 in hscw, to produce CO

1. Oxidation of Fe^{II} in anoxic conditions, at high pH, at 350°C , 25MPa

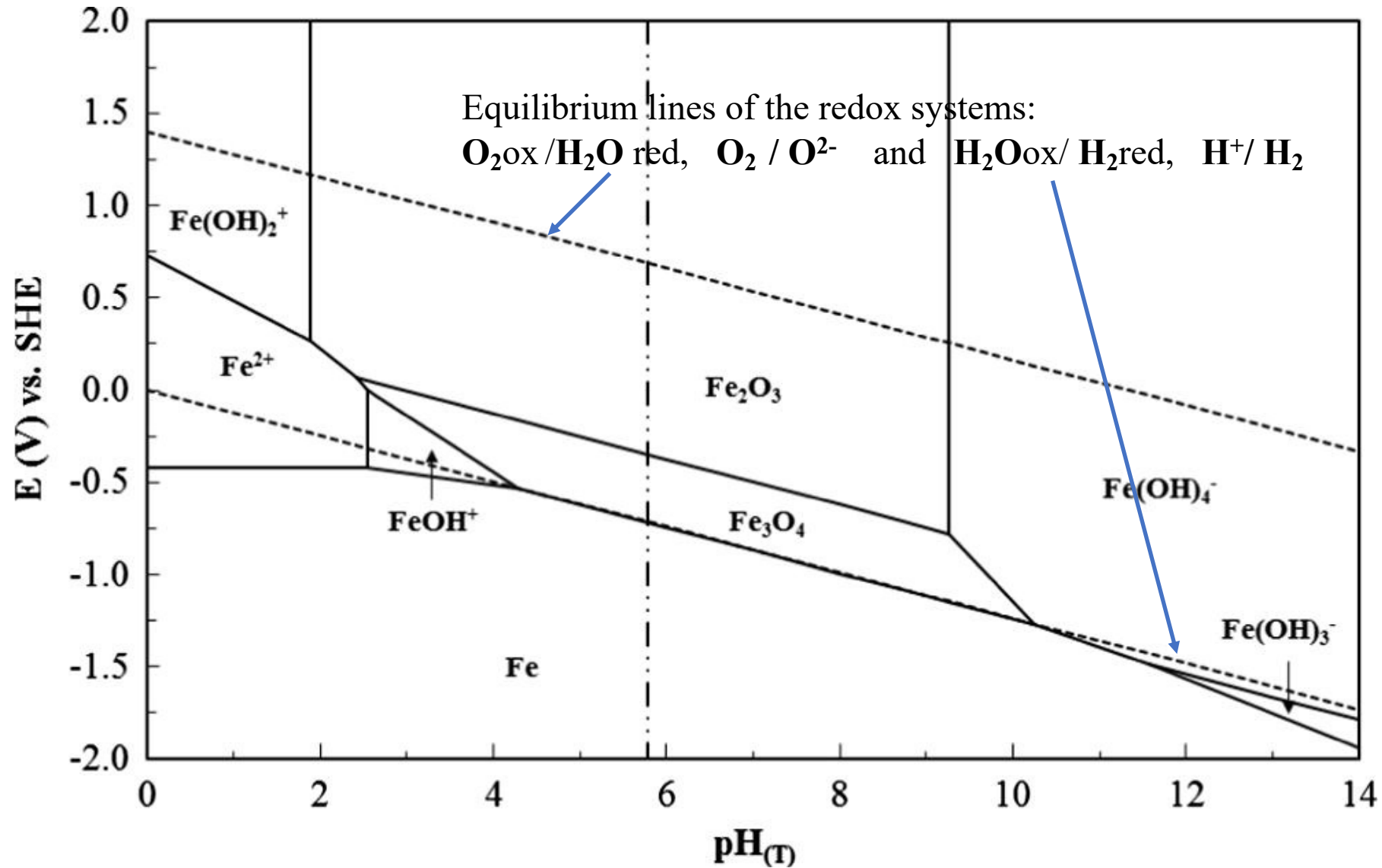


Fig. Pourbaix diagram for iron and water at 350°C , 25 MPa, 10^{-6} mol/kg

Credit: William G. Cook & Robert P. Olive 2012, with permission
Pourbaix diagrams for the iron-water system extended to high & low-supercritical conditions. Corrosion Science 55: 326-331.



2. High Solubility of SiO₂ in high-subcritical water

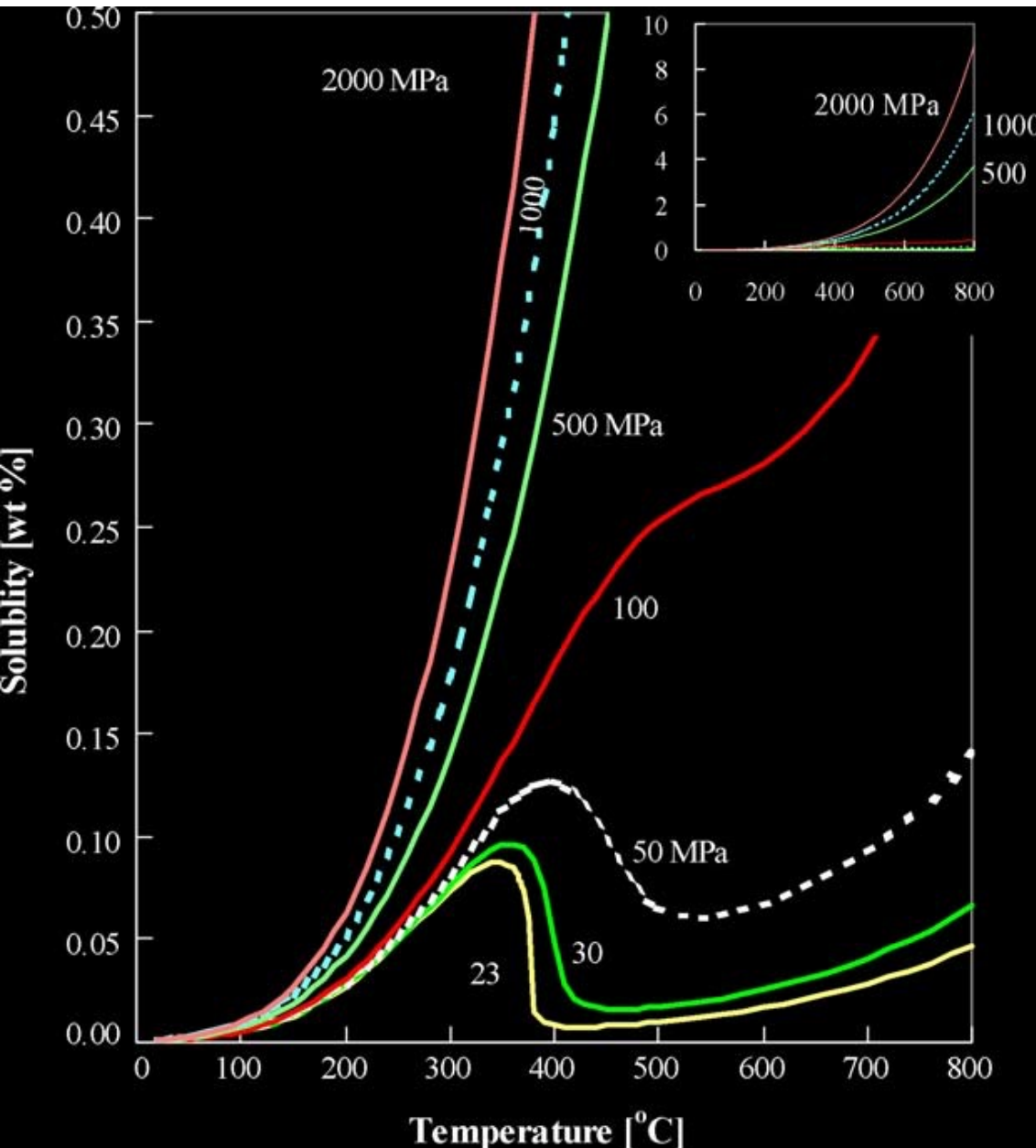


Fig. Solubility of quartz (SiO₂) as a function of T&P

Credit: Richard L. Smith Jr. & Zhen Fang (2011)

Properties and phase equilibria of fluid mixtures as the basis for developing green chemical processes. Fluid Phase Equilibria 302: 65-73.

The solubility of quartz, SiO₂, in water is calculated up to 800°C and 2000 MPa.

« At a constant pressure of **23 MPa**, the SiO₂ solubility increases to a value of 0.087 wt% at **350°C** and drastically decreases to 0.0081 wt% at 450°C »

Everett Shock, Harold Helgeson, Dimitri Sverjensky (1989) Geochim Cosmochim Acta 53:2157-2183

Solubility of quartz: *Kennedy (1950)*

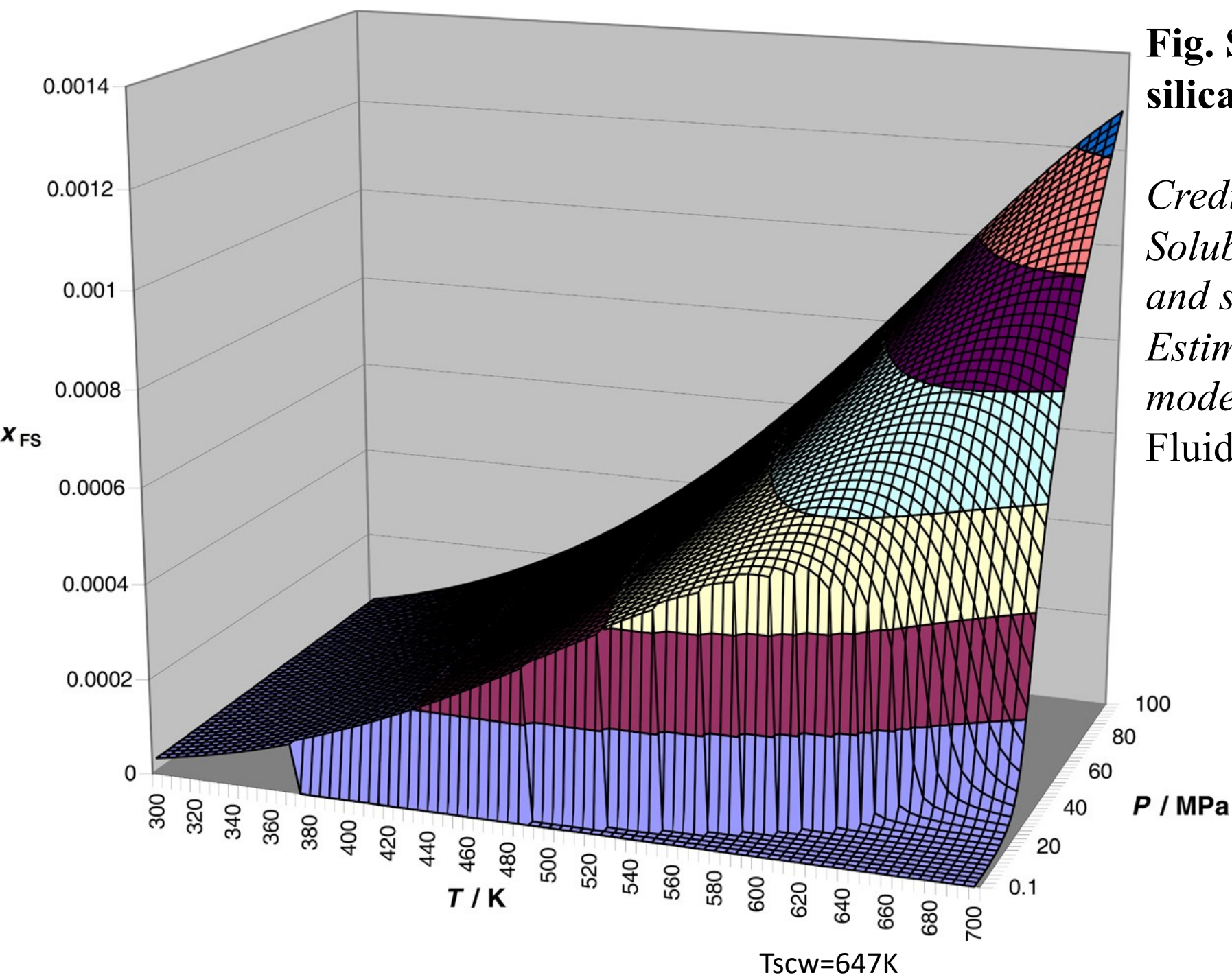


Fig. Solubility of amorphous silica as a function of T&P

Credit: Karasek et al. (2013)
Solubility of fused silica in sub- and supercritical water:
Estimation from a thermodynamic model. The J. of Supercritical Fluids 83: 72-77 with permission

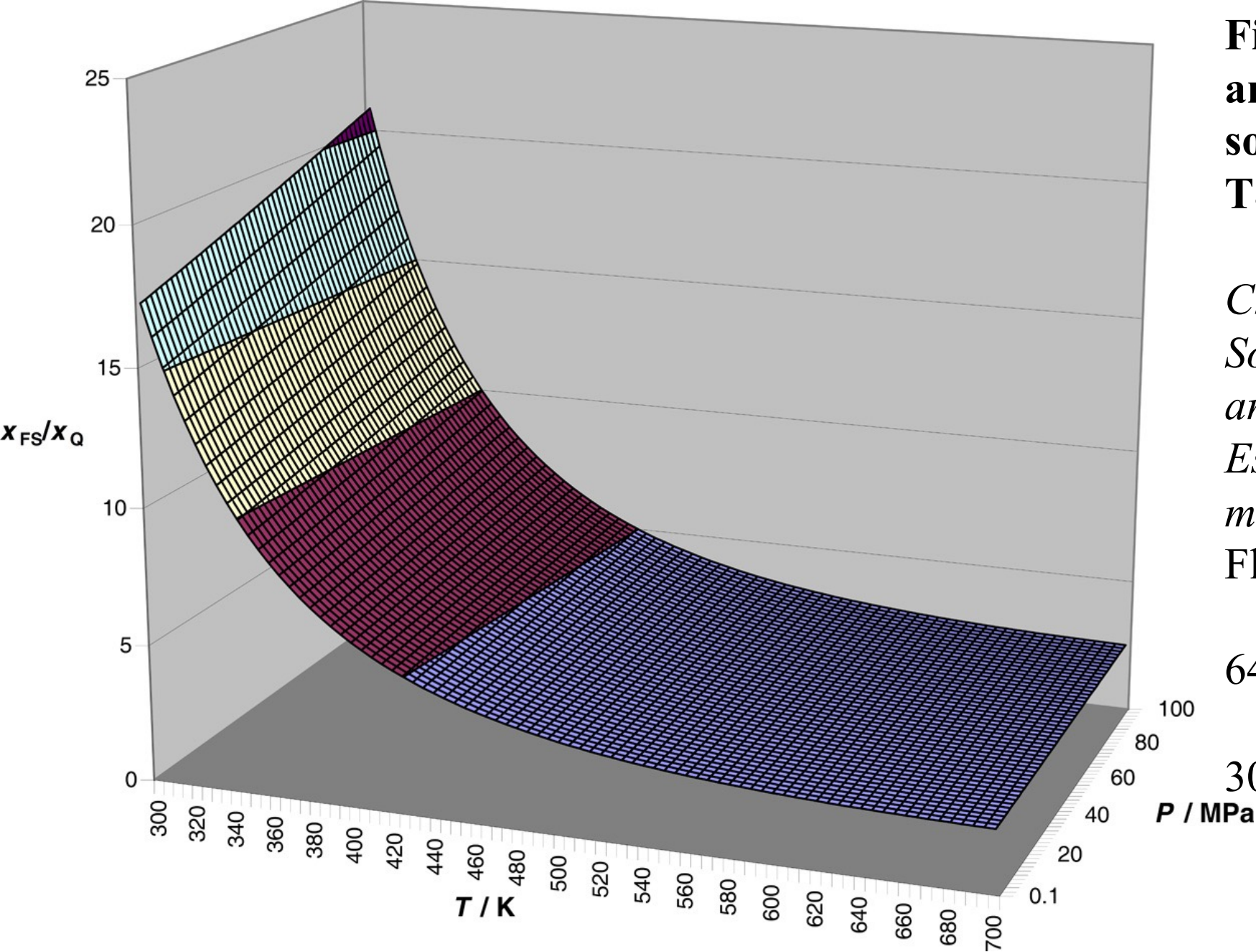


Fig. Ratio of the solubility of amorphous silica to the solubility of quartz at the same T&P

Credit: Karasek et al. (2013)

Solubility of fused silica in sub- and supercritical water:

Estimation from a thermodynamic model. The J. of Supercritical Fluids 83: 72-77 with permission

647K (374°C) & 22.1 MPa:

ratio= ca 2

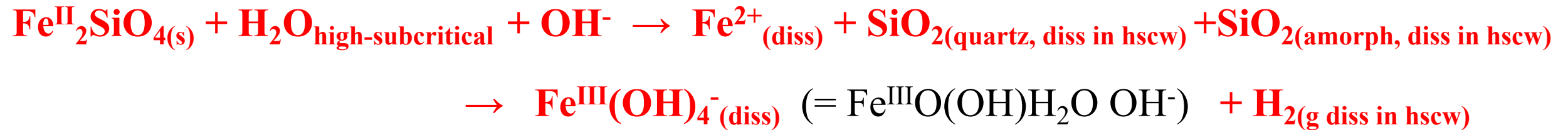
300K(27°C) & 0.1 Mpa (1atm)

ratio= ca 17

3. Interaction between Fe^{II}-silicates with alkaline high-subcritical water for the production of Fe^{III}-oxides, Fe^{III}-silicates and H₂

Formation of ferric oxides

* Dissolution of fayalite and formation of the ferric trihydroxide anion:



* Dehydration of ferric trihydroxide into the ferric oxide hydroxides goethite & lepidocrocite



* Dehydration of the ferric oxide hydroxides into hematite, with possibly ferrihydrite



Formation of ferric silicates

Hydrolysis of fayalite to form Fe^{III}-greenalite



In hscw:

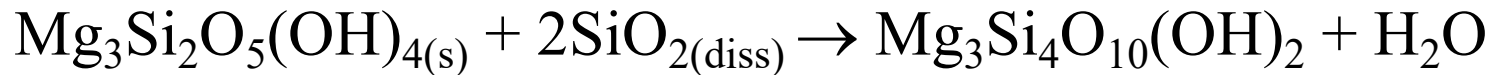


* The iron analog of the serpentine chrysotile is: **greenalite: Fe^{II}₃Si₂O₅(OH)_{4(s)}**



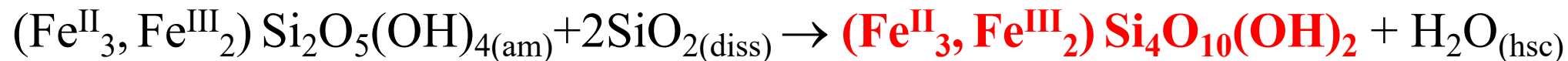
forsterite

chrysotile



talc

* **Dehydration of Fe^{III}-greenalite into Fe^{III}-minnesotaite in a sol. supersaturated in silica**



Fe^{III}-minnesotaite

4. Hydrogenation of CO₂ by H₂ in hscw for production of CO



H_{2(g)} is injected in seawater containing dissolved CO₂ and Fe₃O₄ as catalyst

Prebiotic Chemistry

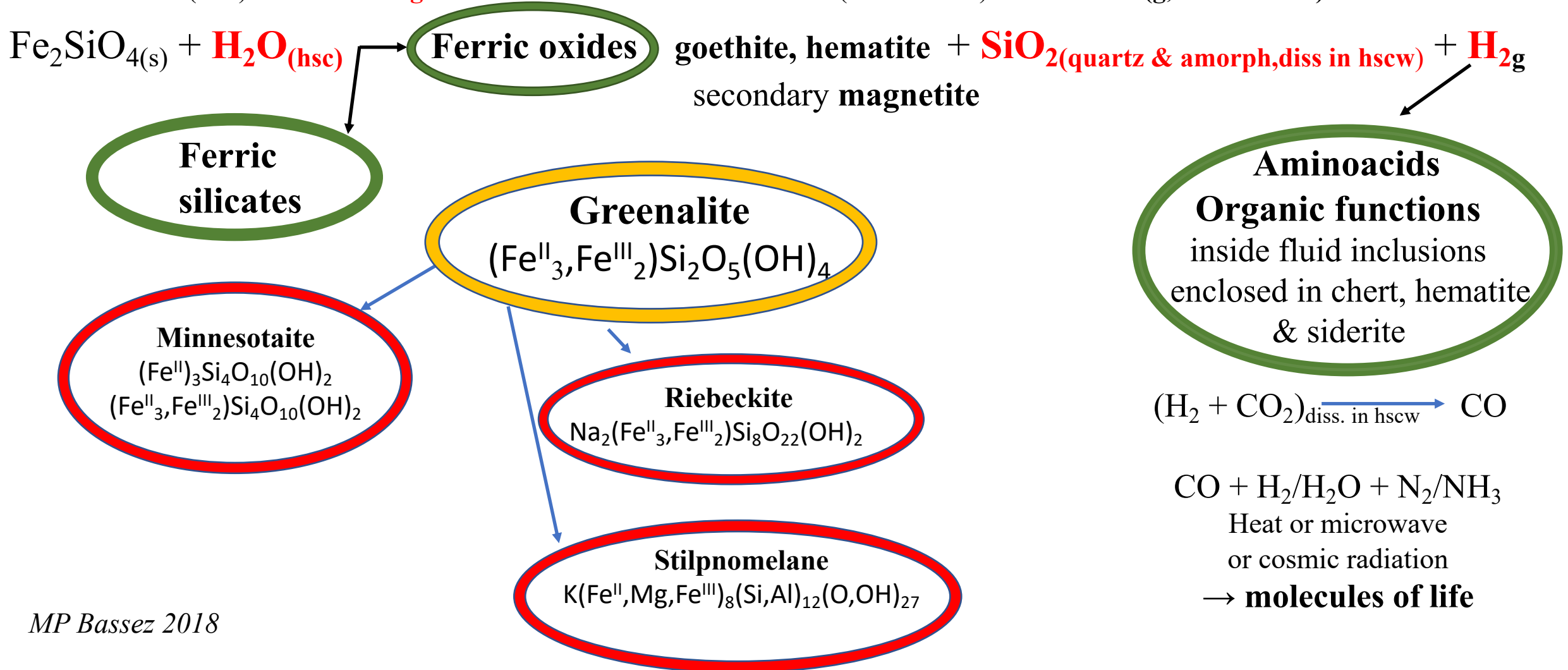
- $\text{CO} + \text{H}_2 + \text{N}_2 \rightarrow \text{CH}_3\text{NH}_2 \quad \text{CH}_3\text{C}\equiv\text{N} \quad \text{CH}_3\text{N}=\text{CH}_2 \quad 300^\circ\text{C} \quad (\textit{Hill \& Nuth 2003})$
precursors of amino acids
- $\text{CO} + \text{H}_2 + \text{NH}_3 \rightarrow \text{amino acids} \quad 374^\circ\text{C} \quad \textit{Pizzarello (2012)}$
- $\text{CO} + \text{H}_2\text{O} + \text{N}_2 \rightarrow \text{amino acids, organic functions} \quad \text{excitation by protons, gamma rays...}$
Kobayashi (1989)...Bassez et al. (2012)

Geobiotropy inside fluid inclusions

The transformation of rocks which contain Fe^{II}-silicates, in alkaline hscw, in the presence of CO₂ & N₂, leads to the synthesis of macromolecules of **life**, at ca 300°-370°C, or 25°C & cosmic radiation

Geobiotropy

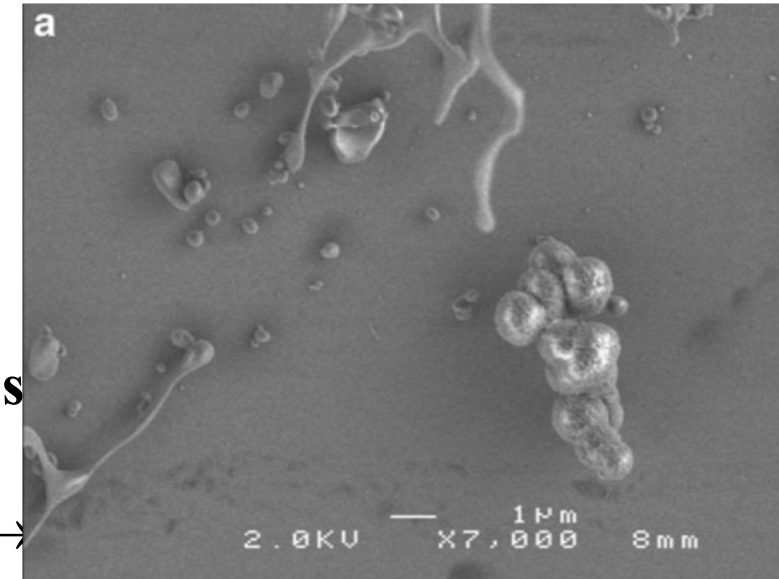
or Follow the water in its high-subcritical state



300°-350°C 10-25MPa

From the numbers emerges the form

- 1. $\text{Fe}^{\text{II}}(\text{OH})_3^-$ diss + $\text{H}_2\text{O}_{\text{high-subcritical}}$ → $\text{Fe}^{\text{III}}(\text{OH})_4^-$ diss in hscw + $1/2\text{H}_2$ diss in hscw** **pH~9.5-14**
no Fe^{III} above Tc (374°C, 22.1MPa) (Bassez 2013)
- 2. 3SiO_2 quartz → SiO_2 diss quartz + ~ 2 SiO_2 amorph** (Shock et al.1989) (Smith&Fang 2011) (Karasek et al. 2013)
no SiO_2 diss above Tc
- 3. CO_2 diss + H_2 diss → CO diss + CH_4 diss** 250°-300°C, 25MPa, catalyst: Fe_3O_4 (Fu & Seyfried 2009)
- 4. $\text{Fe}_2\text{SiO}_4(\text{s})$ + $\text{H}_2\text{O}_{(\text{hsc})}$ → SiO_2 diss Q + SiO_2 amorph silica**
→ **Ferric oxides:** goethite, hematite + H_2 (diss in hscw)
→ **Ferric silicates:** greenalite $(\text{Fe}^{\text{II}}_3, \text{Fe}^{\text{III}}_2)\text{Si}_2\text{O}_5(\text{OH})_4$
minnesotaite $(\text{Fe}^{\text{II}}_3, \text{Fe}^{\text{III}}_2)\text{Si}_4\text{O}_{10}(\text{OH})_2$
 $(\text{Fe}^{\text{II}}_3, \text{Fe}^{\text{III}}_2)\text{Si}_2\text{O}_5(\text{OH})_4 + \text{SiO}_{2(\text{diss})} \rightarrow (\text{Fe}^{\text{II}}_3, \text{Fe}^{\text{III}}_2)\text{Si}_4\text{O}_{10}(\text{OH})_2 + \text{H}_2\text{O}$
→ **H_2**
 $\text{H}_2 + \text{CO}_2 \rightarrow \text{CO}$; $\text{CO}_g + \text{H}_2\text{O}_g + \text{N}_{2g} + \gamma, \text{ protons..} \rightarrow \text{organic functional groups}$
(Kobayashi et al.1990, room T) (Takano et al. 2004, 765°C,10s)
(Kurihara et al. 2012, 300°C, 25MPa, 2min) (Bassez, Takano, Kobayashi 2012) →



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